

The role of anisotropy and heterogeneity in Geothermal systems in meta-sedimentary rocks

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Content

Havelange demo-site presentation

What is anisotropy and heterogeneity?

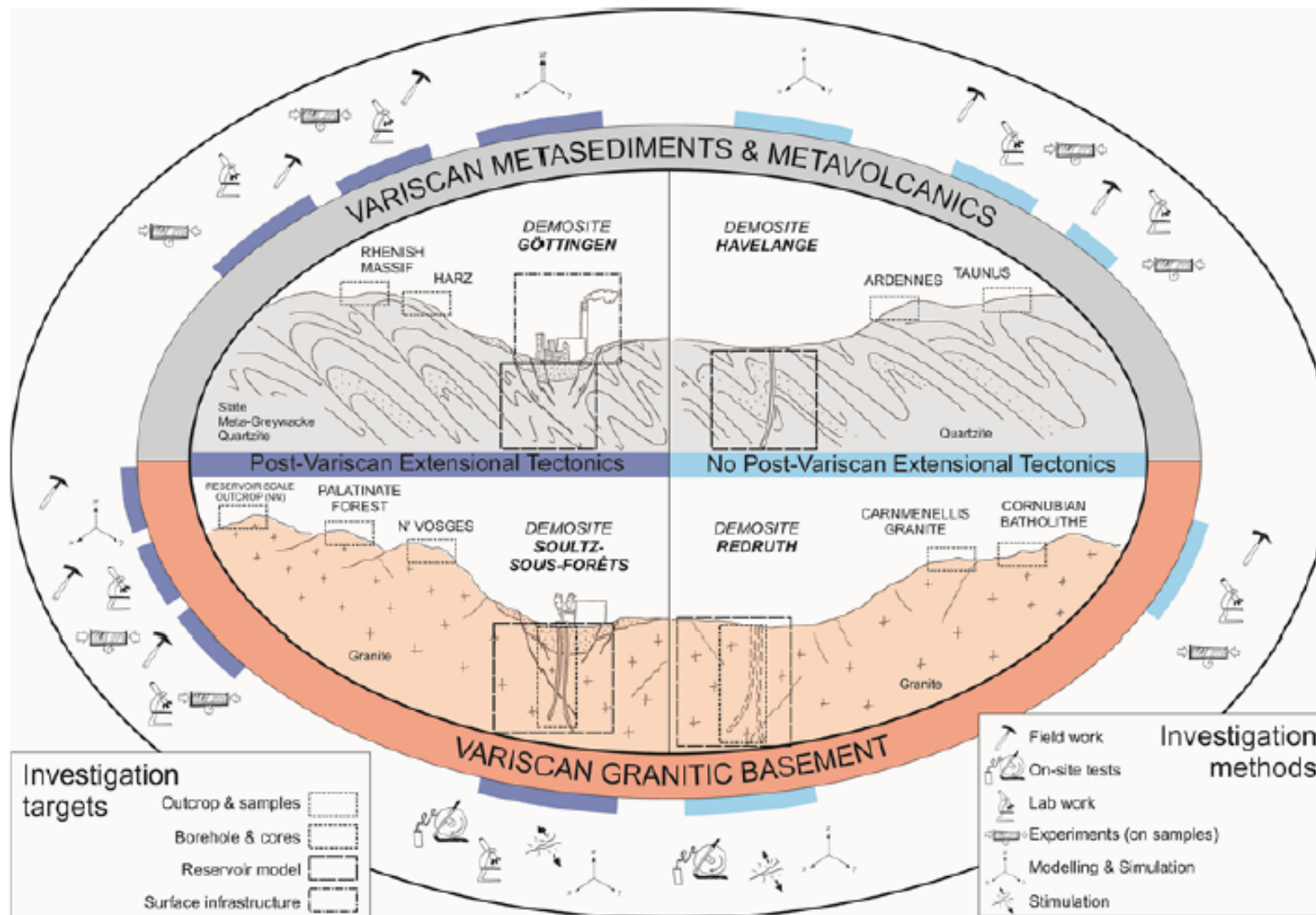
Anisotropy in Geology

Study case 1: thermal conductivity

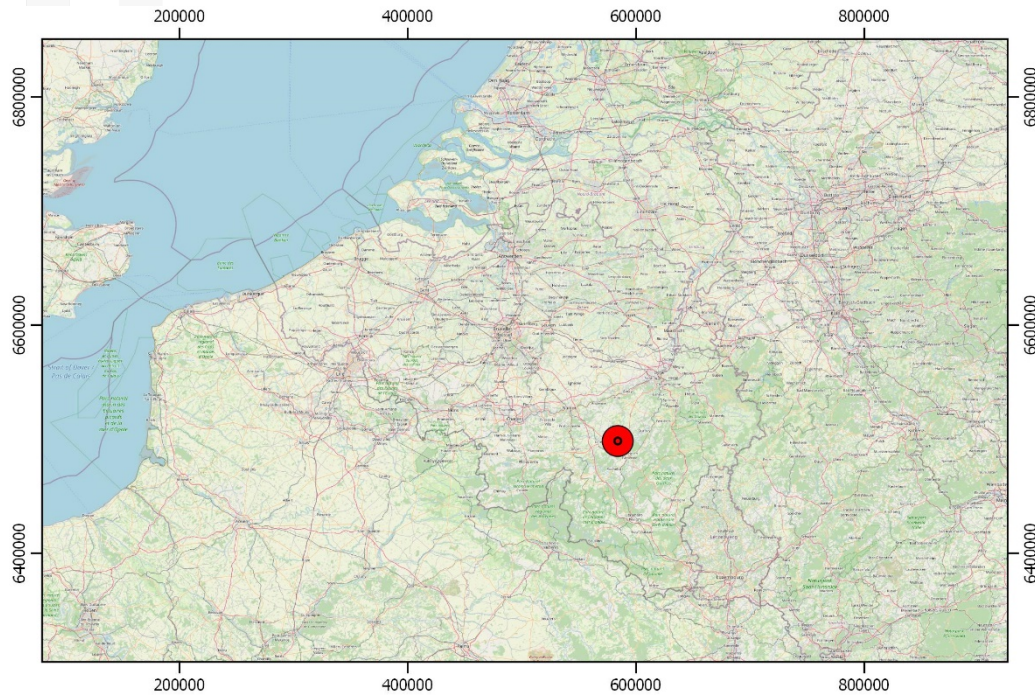
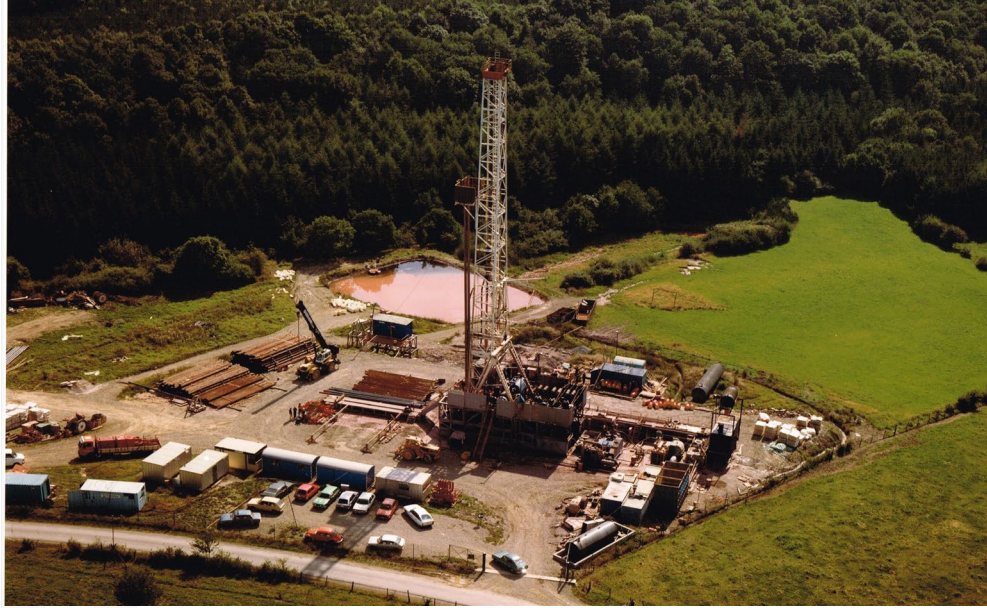
Study case 2: rock fracturation in metasedimentary formations

Take-away message

Havelange demo-site in the MEET project



Havelange borehole ID-card



Drilled from 1981->1984 near the village of Netinne (Belgium)

Reached a maximum depth of 5648 m (MD)

Aim: gas exploration well targeting hidden Upper Carboniferous coal measures under the main detachment level (Midi Thrust Fault)

Encountered stratigraphy from Upper- to Lower-Devonian formations

Located in the Dinant Synclinorium: a sub-unit of the Ardenne Allochthon part of the Rhenohercynian fold-and-thrust belt (Variscan Orogeny)

Located about 23 km South of the Variscan Front

Geothermal targets: Lower Devonian quartzite units observed ~4.3 -> 5.3 km

Recorded temperature (down hole): 126°C

Havelange available information



Cuttings samples (one sample every 1 to 5 m) + cores at shallow depth (Famennian shales) and at great depth (quartzite)

Logs (paper-format) -> during the MEET project: GR, Dipmeter, Sonic, Caliper were digitized

8 seismic lines were shot in 1978

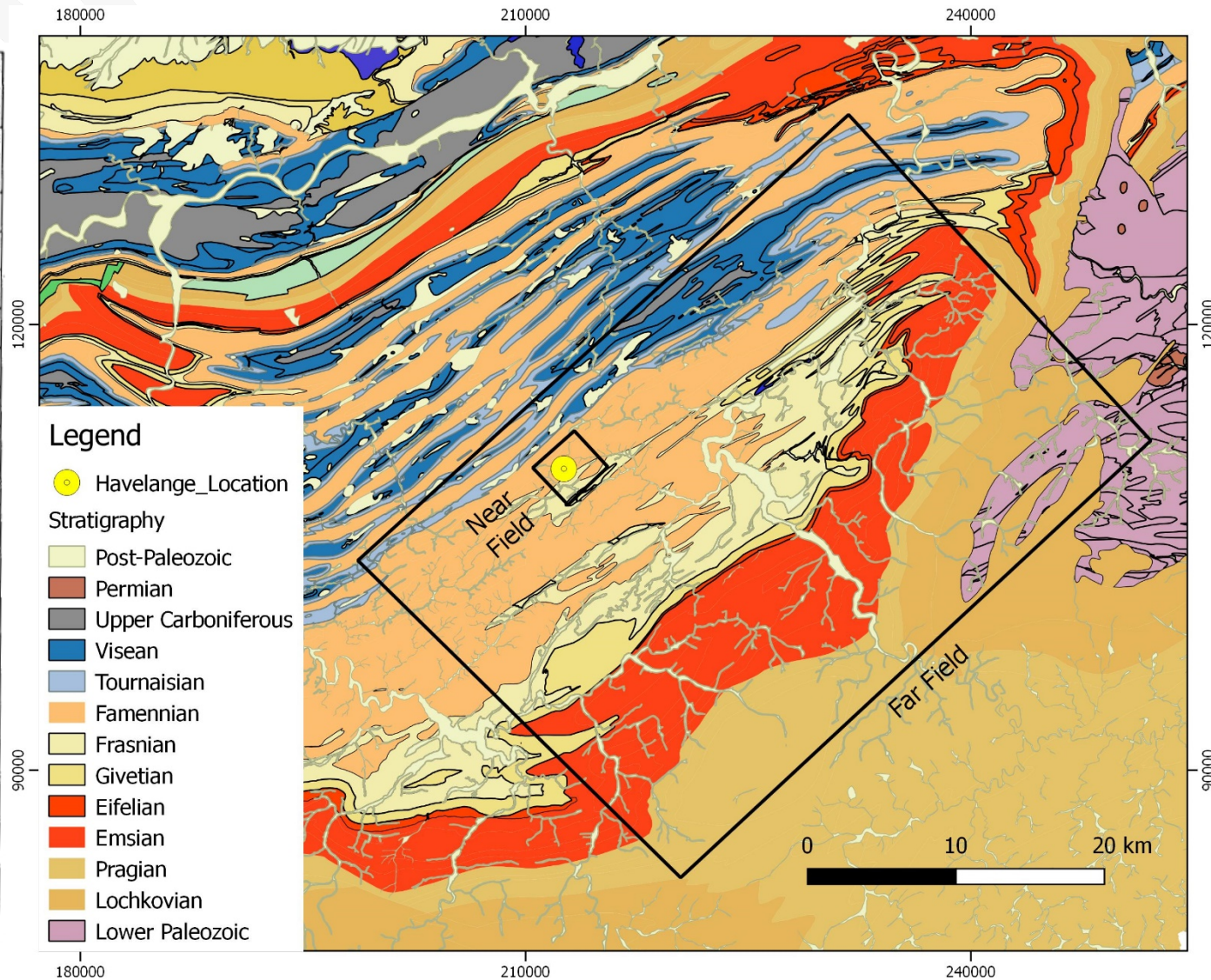
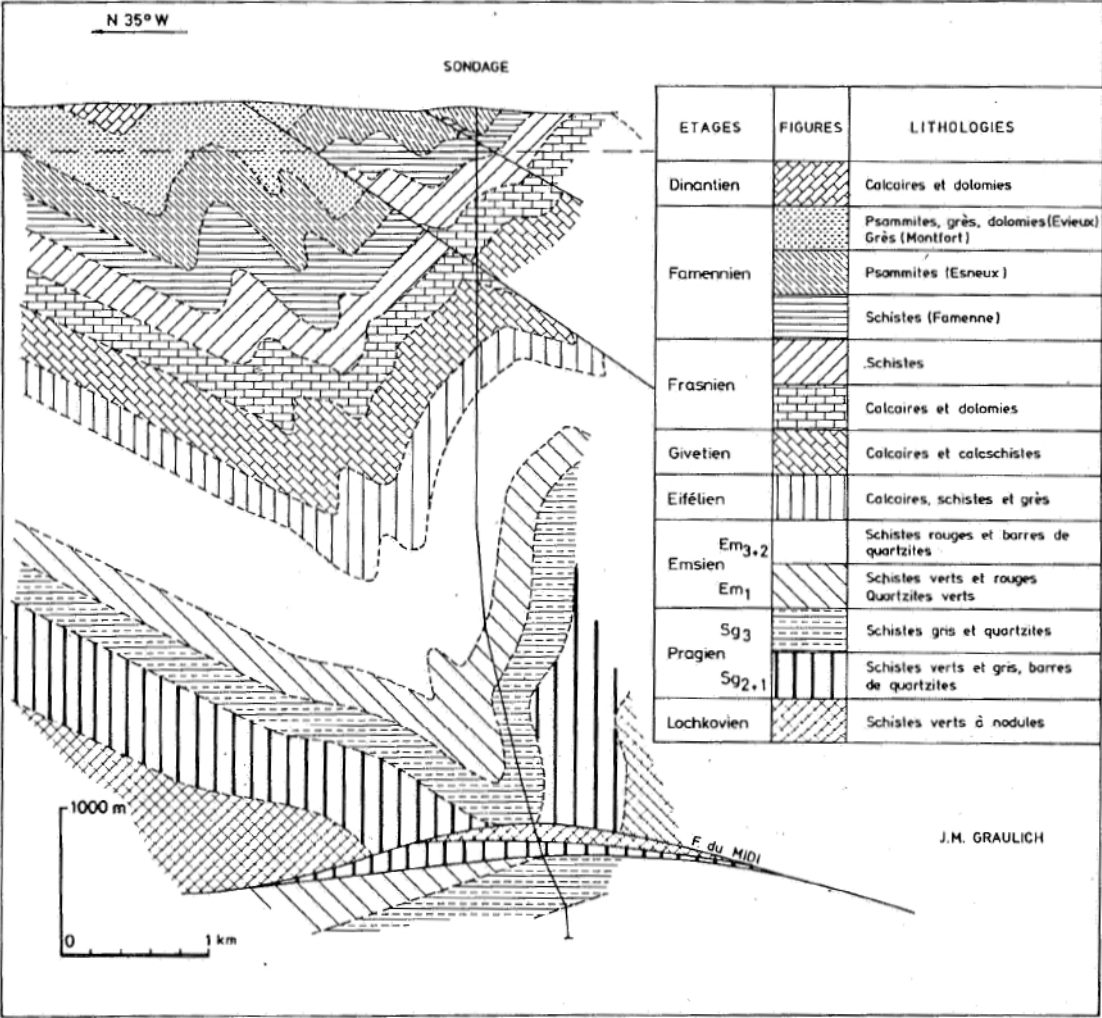
During the MEET project, additional information were collected:

- Restudy of the mineralogical composition of cutting and core samples;
- Restudy of daily drilling report (mud lost analyses);

In the near-field: main activities were the acquisition of drone image during the drought period of Summer 2018 and sampling of spring water

In the far-field (analogue field): main activities were study of outcrops, collect rock samples for rock mechanical tests and sampling of spring water

Havelange analog field and field works



What is anisotropy?

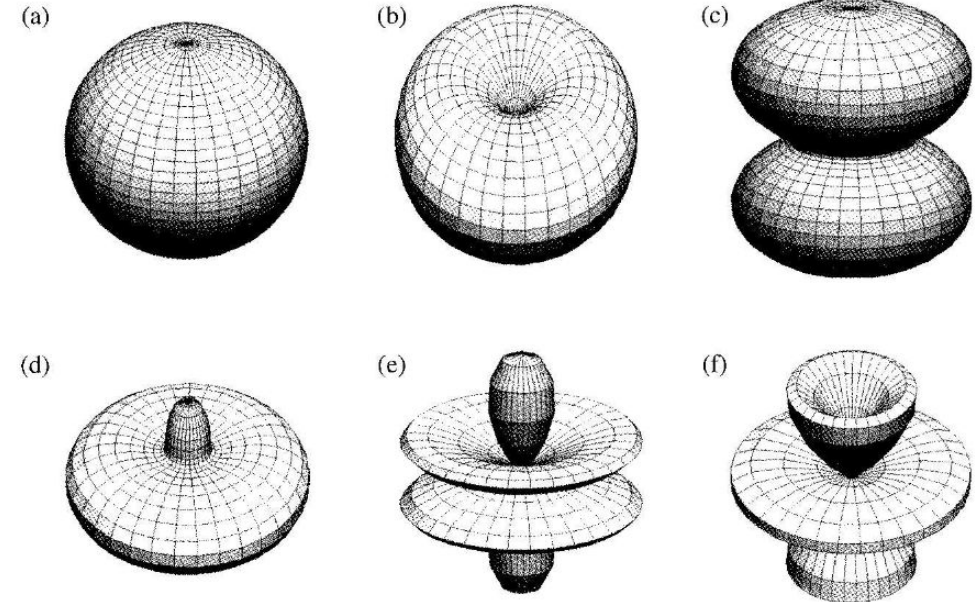
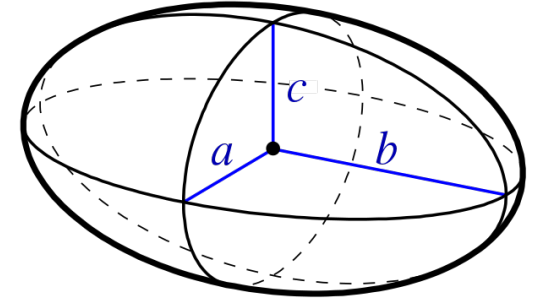
The definition of Wikipedia: Anisotropy is the property of a material which allows it to change or assume different properties in different directions

Anisotropy >< Isotropy

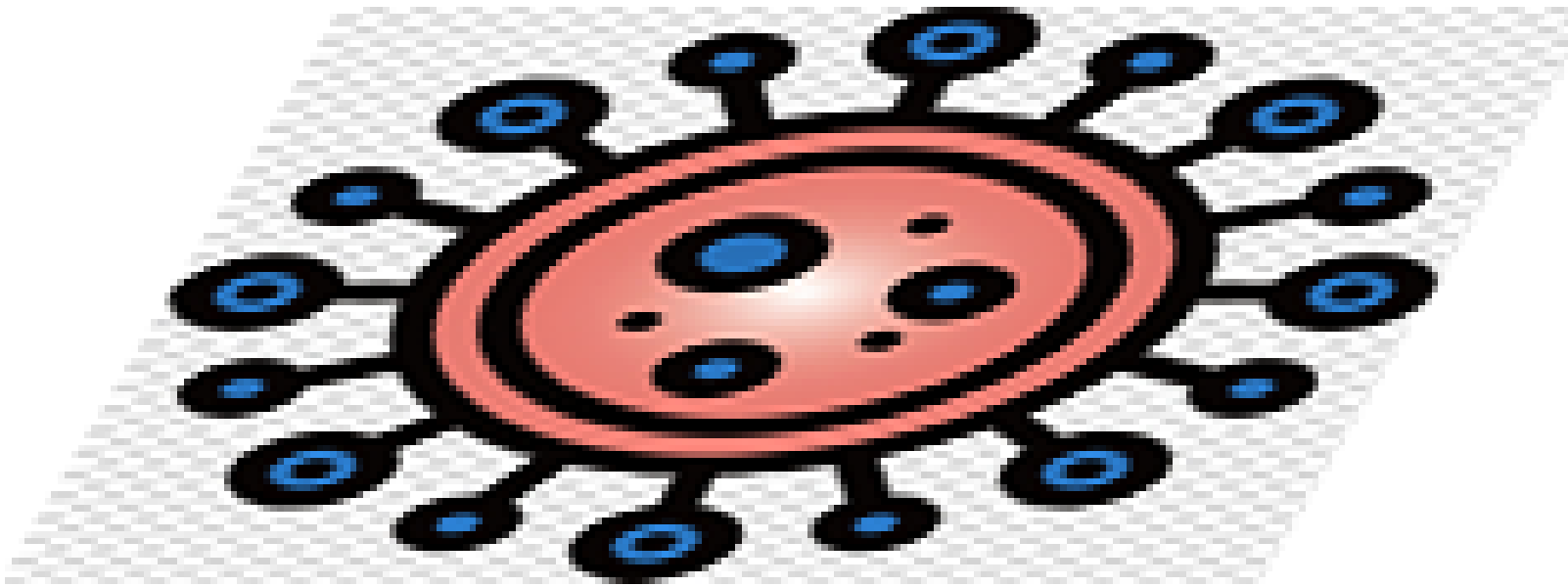
For the simple cases: anisotropy in 2D can be defined as a ratio: $\text{Param_in_max_direction} / \text{Param_in_min_direction}$ (@90°) and for those cases the anisotropy is represented by an ellipse(2D)/ellipsoid (3D).

The isotropy is represented by a circle(2D)/sphere (3D).

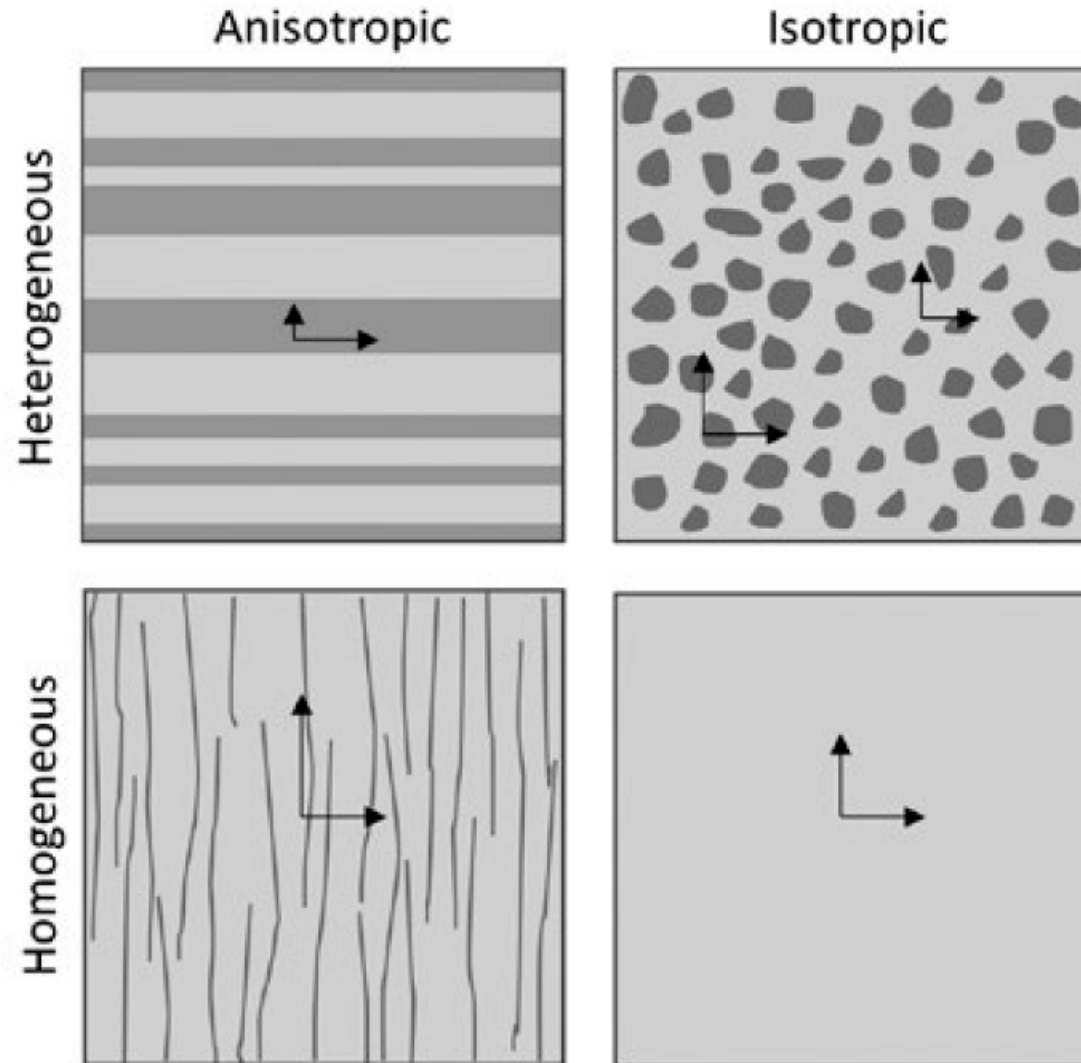
In other cases the anisotropy representation is more complex.



Or even has a weird shape ...



Anisotropy vs heterogeneity?



Lynn (2018)

Anisotropy in Geology

Anisotropy in Geology is present in numerous geological settings, at **different scales** regarding several properties.

Some examples of anisotropy in rocks: mineral plasticity, elasticity, fabric, sonic, electrical conductivity, thermal conductivity & expansion, magnetism, permeability ...

Anisotropy is also strongly **link to a scale**: from a single crystal (plasticity, electrical conductivity), to hand-specimen or an outcrop for a field geologist or even at the scale of geological sequences or formations

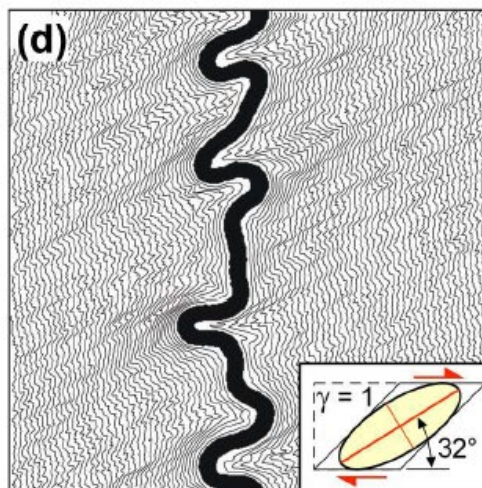
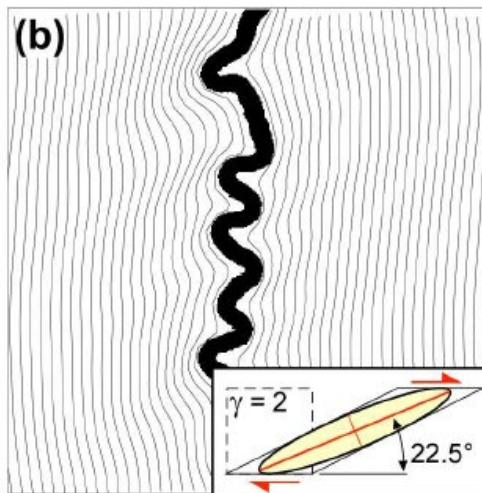
The origin of anisotropy in a rock happens at different moment in the Geological cycle: intrinsic or primary anisotropy during sedimentary rock deposit and diagenesis to secondary anisotropy developing during tectonic processes (e.g. cleavage development related to stress)

Is it important to take in account anisotropy during modelling?

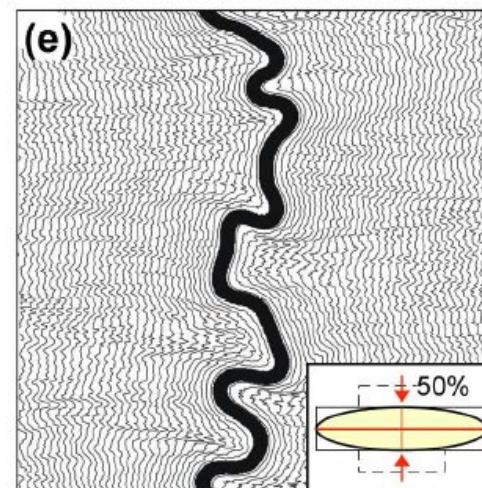
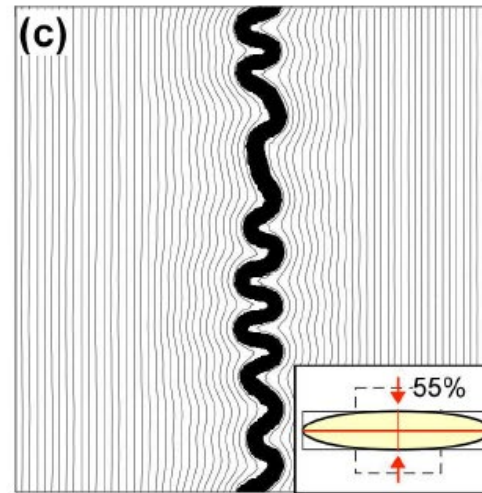


Ran et al. (2018)

Crenulation
cleavage



Simple shear



Pure shear

Isotropic rheologies
Folding seeds =
Variation of single layer
thickness

« Anisotropic »
rheologies taking into
account mechanical
anisotropy of the
matrix.
Folding result from
random variations in
lattice orientation

Study case 1: thermal conductivity anisotropy

Reminder:

Heat transfer mechanisms: conduction, convection and radiation

Let's consider a 1D-approach of Fourier's Law of Heat Conduction:

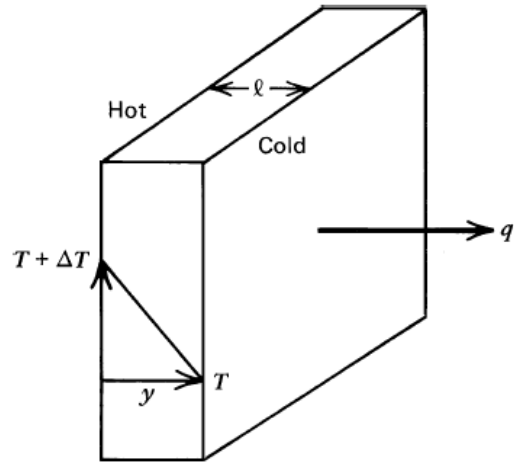
q is the heat flux

$$q = -k \frac{\Delta T}{l} \quad [\text{W/m}^2]$$

k is the coefficient of thermal conductivity [$\text{Wm}^{-1}\text{K}^{-1}$]

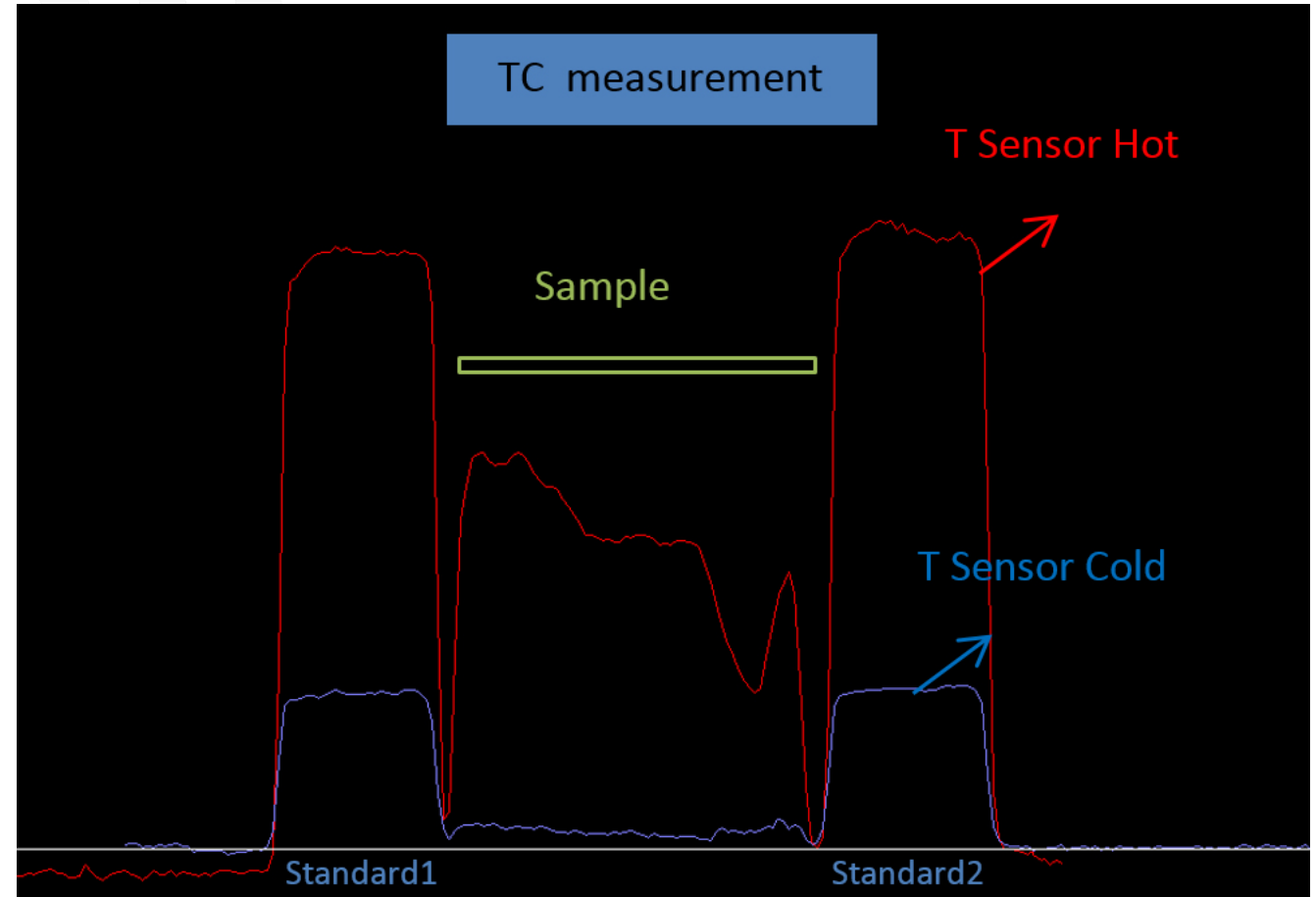
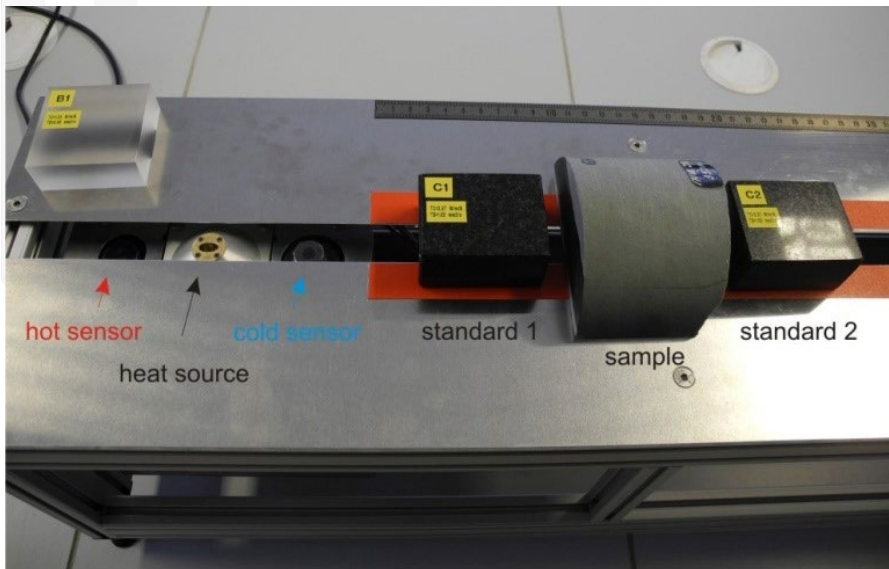
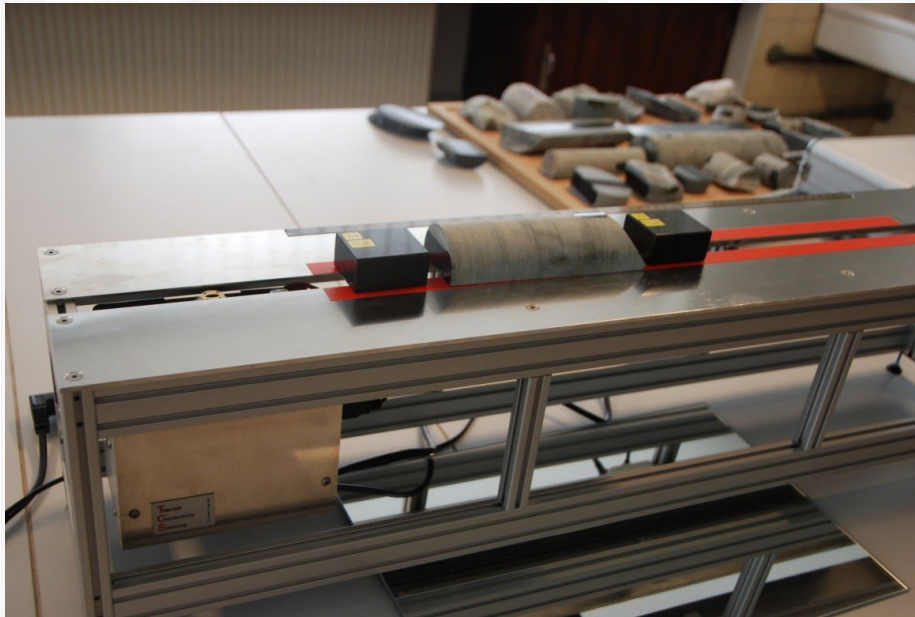
k for rocks ranges from ~ 1.2 (shale) to ~ 5 (dolomite, quartzite)

k values can be measured in lab or directly in borehole conditions



(Turcotte & Schubert, 2002)

Thermal Conductivity Scanning – TCS



(Popov et al., 1999, 2012 ; Hartmann, 2005)

Thermal conductivity anisotropy

Regarding metasedimentary formations:

Laminated rocks conduct heat preferentially parallel to bedding planes

k values measurement are conducted on surface perpendicular to bedding according for at least 2 directions: parallel to the bedding (k_{par}) and perpendicular to the bedding (k_{perp} , usually = k_{min})

We can define the anisotropy as the ratio $k_{\text{par}}/k_{\text{perp}}$. According to Davis et al. (2007), anisotropy values range between 0.8 and 2.1.

For a given angle we can derive an apparent thermal conductivity (k_{app})

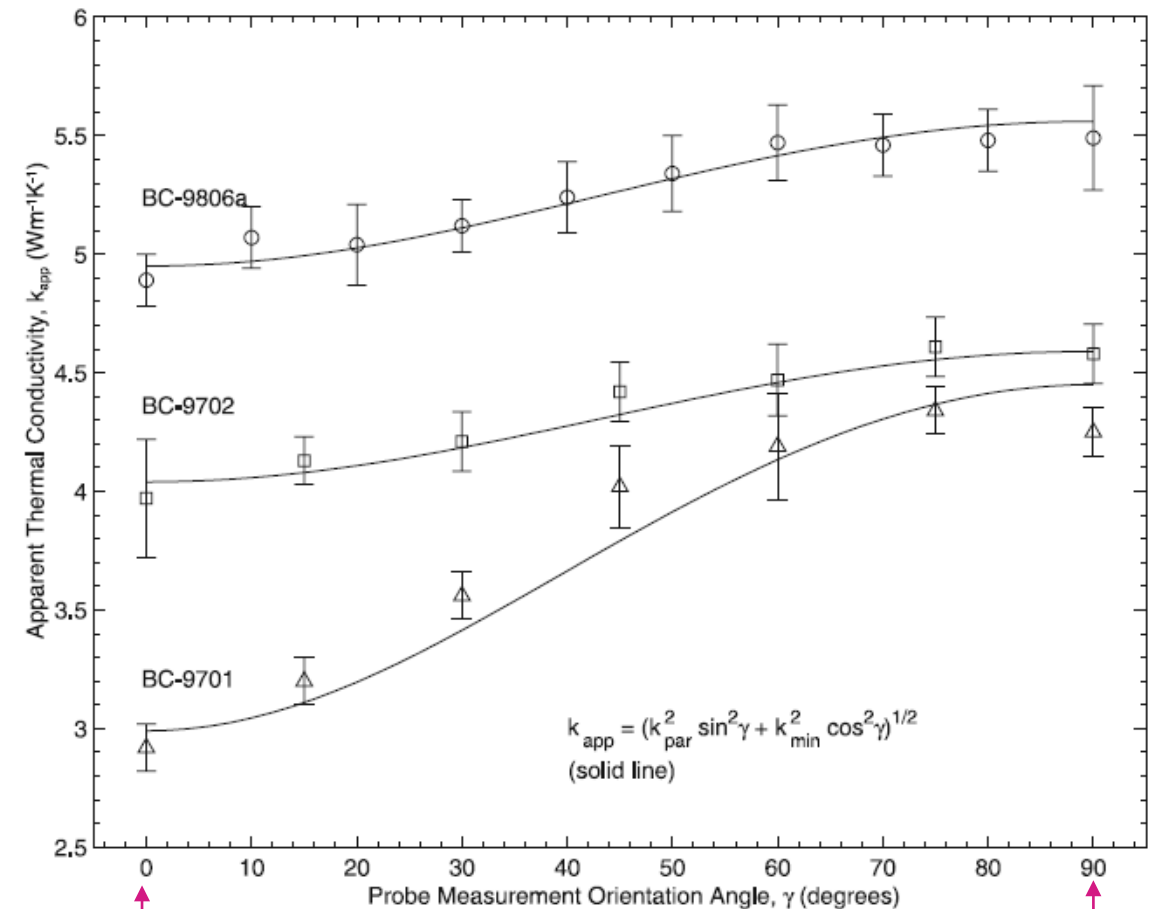


Figure 2. Apparent thermal conductivity as a function of probe orientation for samples BC-9701, BC-9702 and BC-9806a. BC-9806a has been offset by a value of 1 $\text{Wm}^{-1}\text{K}^{-1}$ to differentiate it from the other samples. Bars indicate the standard deviation of six or more measurements during one sample determination.

Measurement \perp to S0

Measurement \parallel to S0

Exercise 1 – Impact of anisotropy of q

If we consider the Fourier's Law in 1D applied on the measurements conducted by Davis on sample BC-9701.

What is the computed k_{app} if the probe measurement orientation angle 40° ?

What is the impact on q -value if the heat flux (Y) occurs perpendicular to the rock fabric with respect to the flux occurring // to this fabric?

Hints: you can consider for instance a difference of 3° between the hot and cold faces of the slab and its thickness is 100 m

$$q = -k \frac{\Delta T}{l} \quad [\text{W/m}^2]$$

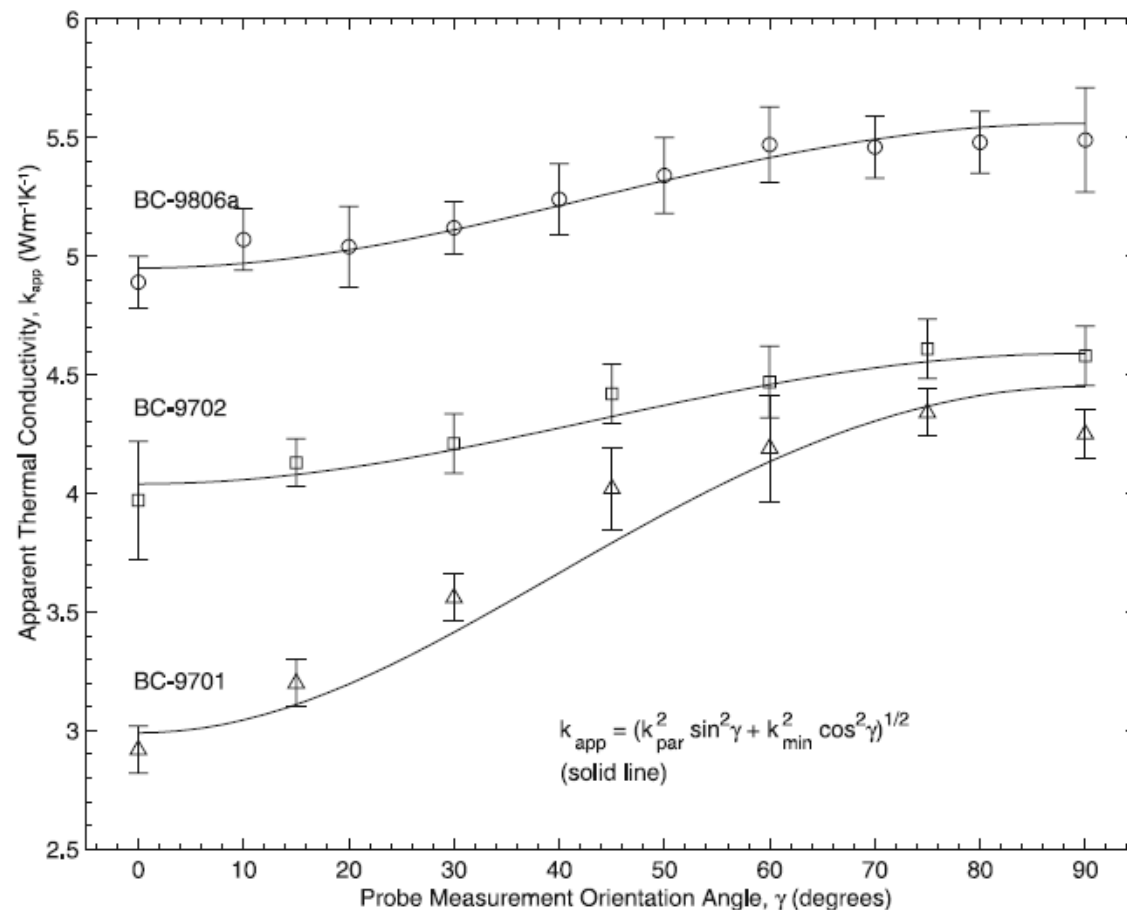


Figure 2. Apparent thermal conductivity as a function of probe orientation for samples BC-9701, BC-9702 and BC-9806a. BC-9806a has been offset by a value of $1 \text{ Wm}^{-1}\text{K}^{-1}$ to differentiate it from the other samples. Bars indicate the standard deviation of six or more measurements during one sample determination.

Exercise 1 – Impact of anisotropy of q

If we consider the Fourier's Law in 1D applied on the measurements conducted by Davis on sample BC-9701.

What is the computed k_{app} if the probe measurement orientation angle 40° ?

$k_{app} = 3.67 \text{ W m}^{-1} \text{ K}^{-1}$

What is the impact on q -value if the heat flux (Y) occurs perpendicular to the rock fabric with respect to the flux occurring // to this fabric?

Hints: you can consider for instance a difference of 3° between the hot and cold faces of the slab and its thickness is 100 m

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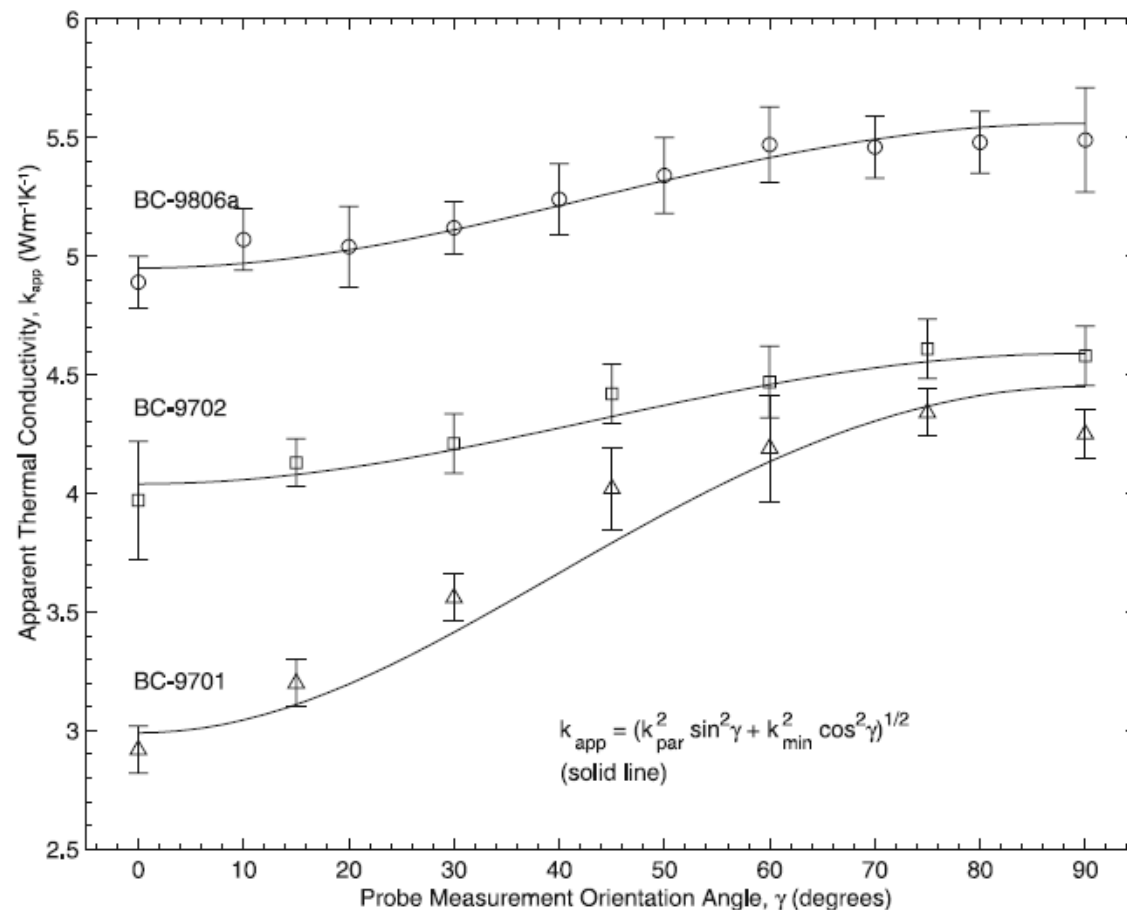


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$$q_{//} = 4.45 \cdot 0.03 = 0.134 \text{ Wm}^{-2}$$

$$q_{\perp} = 3.00 \cdot 0.03 = 0.090 \text{ Wm}^{-2}$$

Reduction of $\sim 33\%$ of q

$$q = -k \frac{\Delta T}{l} \quad [\text{W/m}^2]$$

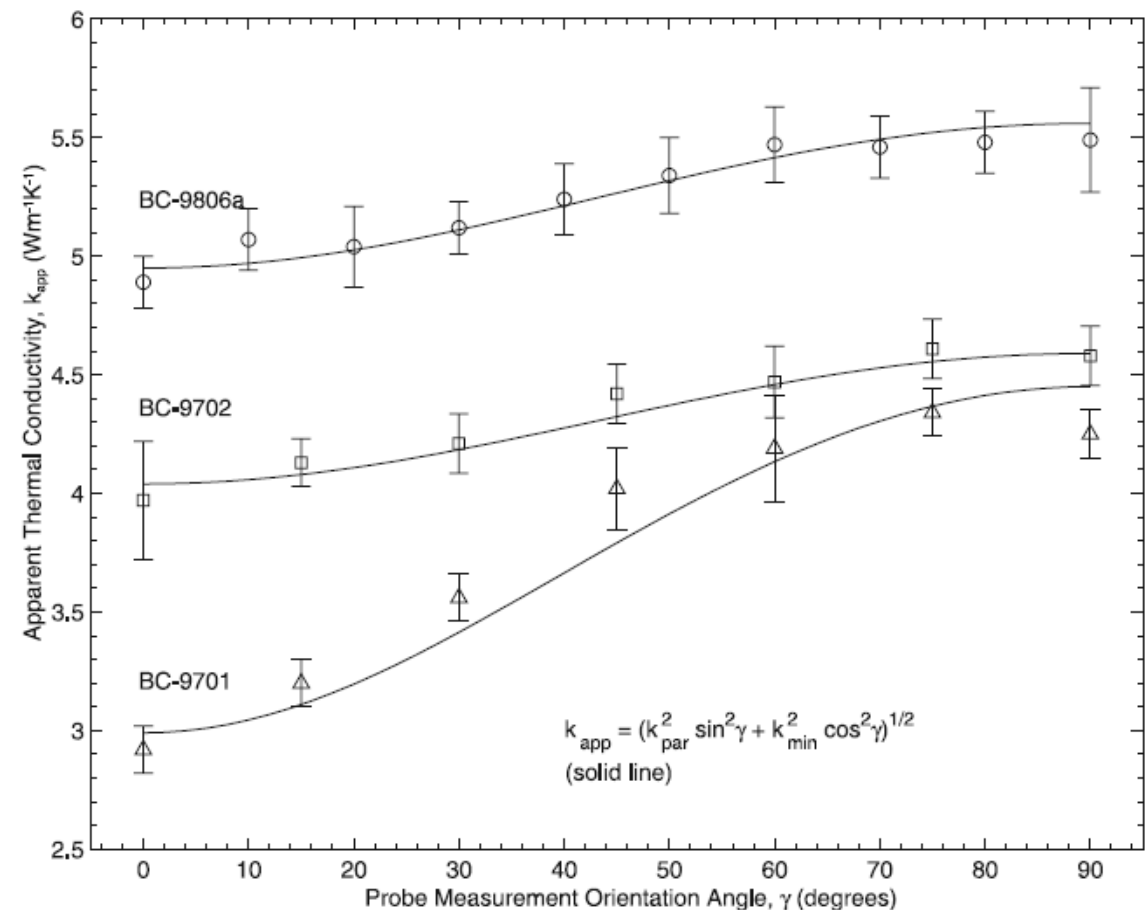


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Additional remarks/thoughts

- The impact of thermal conductivity anisotropy is significant mainly for strongly laminated rocks (e.g. shale, slate). For non-fracturate granite: thermal conductivity \rightarrow isotropy
- For highly anisotropic materials it would be better to provide 2 k-values: one parallel to the bedding and one perpendicular
- The presented approach here is oversimplified (1D) and we have consider only the heat transfer of conduction, other factors or processes should be taken into account are: the convection \rightarrow permeability anisotropic, but also the heat generation

Study case 2: Rock fracturation in cases of heterogenous sequences

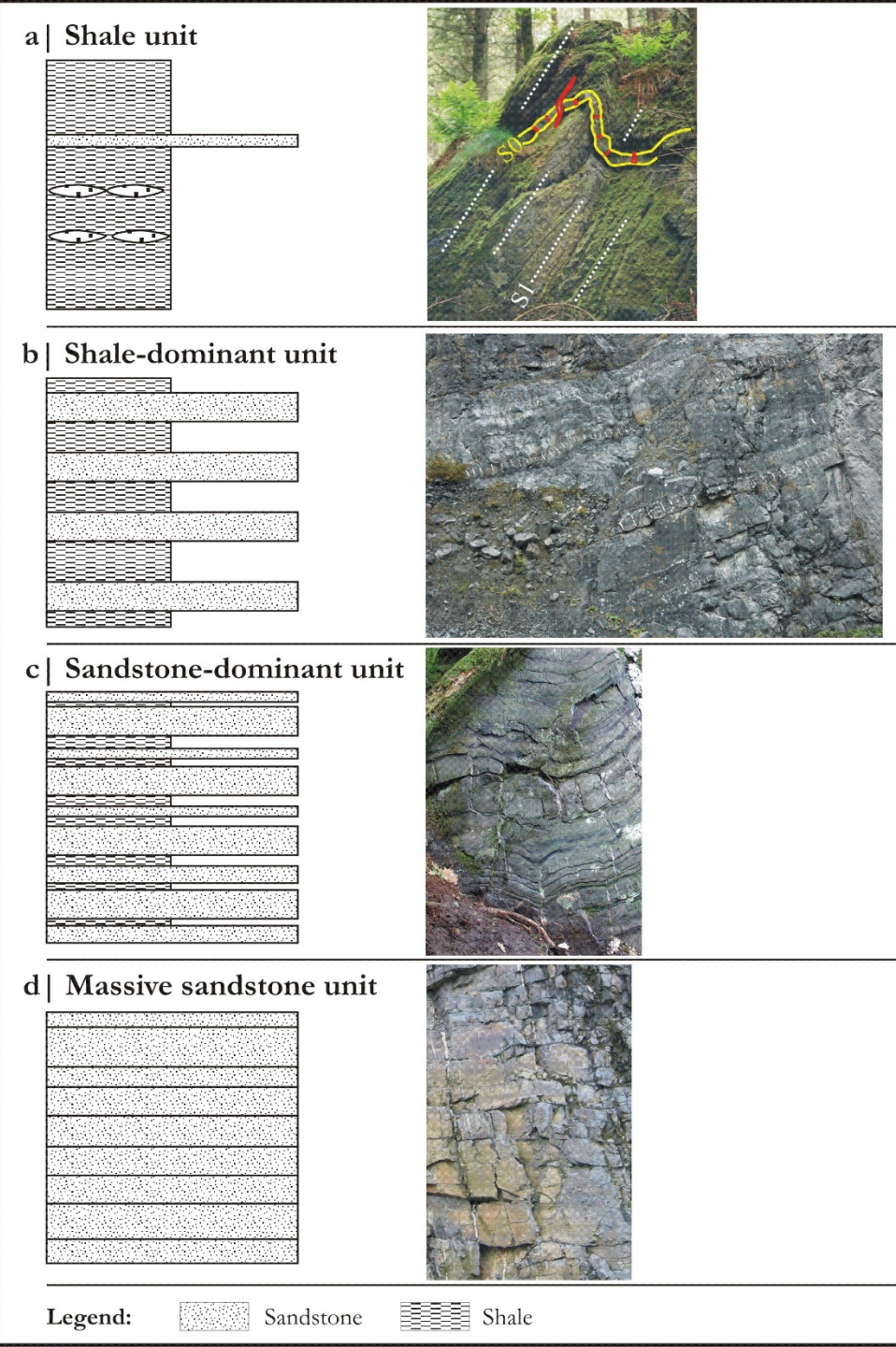
Presented cases come from the Belgian Ardenne

Lithologies: sandstone, quartzite and shale/slate layers

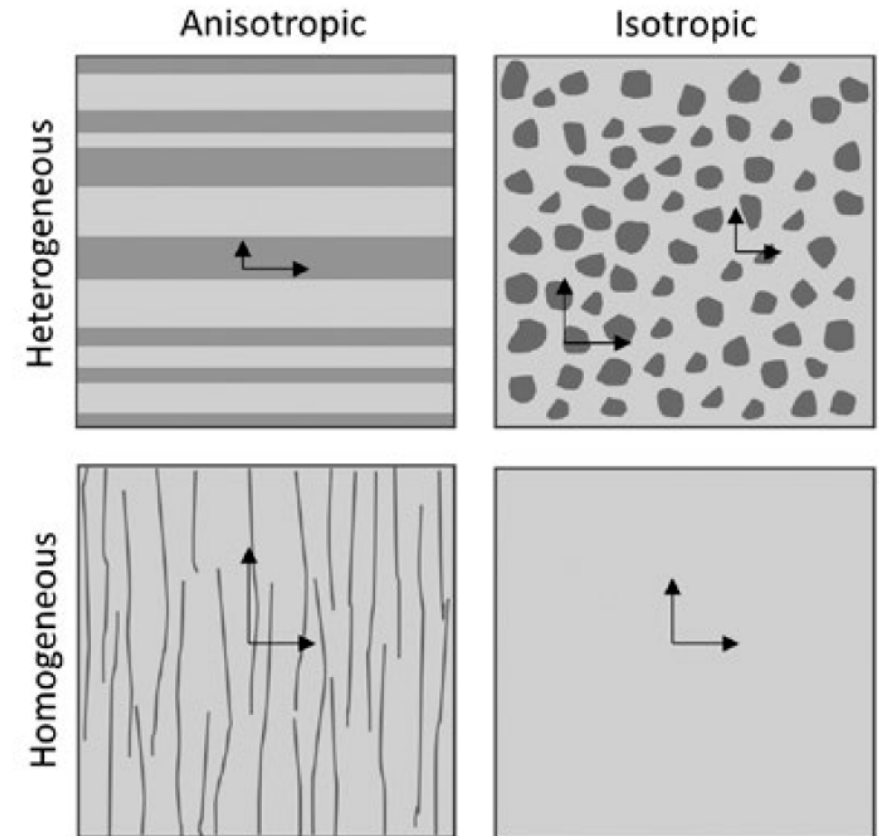
Rock sequences underwent first an extensional period during the metamorphism peak (basin in mature state) followed by a shortening period during the Variscan orogeny (330-300 Ma)

During the extensional period, rock formations underwent fracturation with the development of quartz veins.

Metamorphism grade: green-schist facies



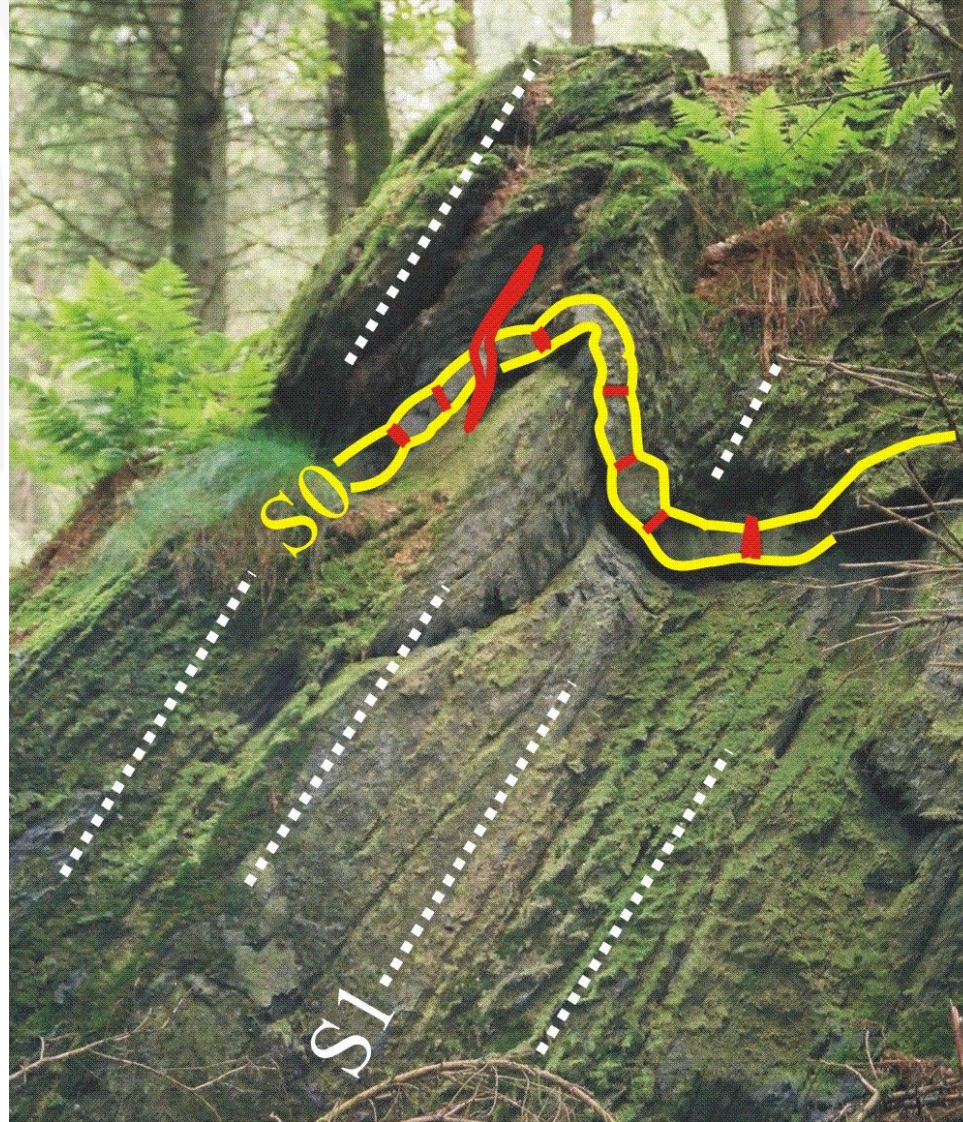
Fractures in a massive sandstone unit



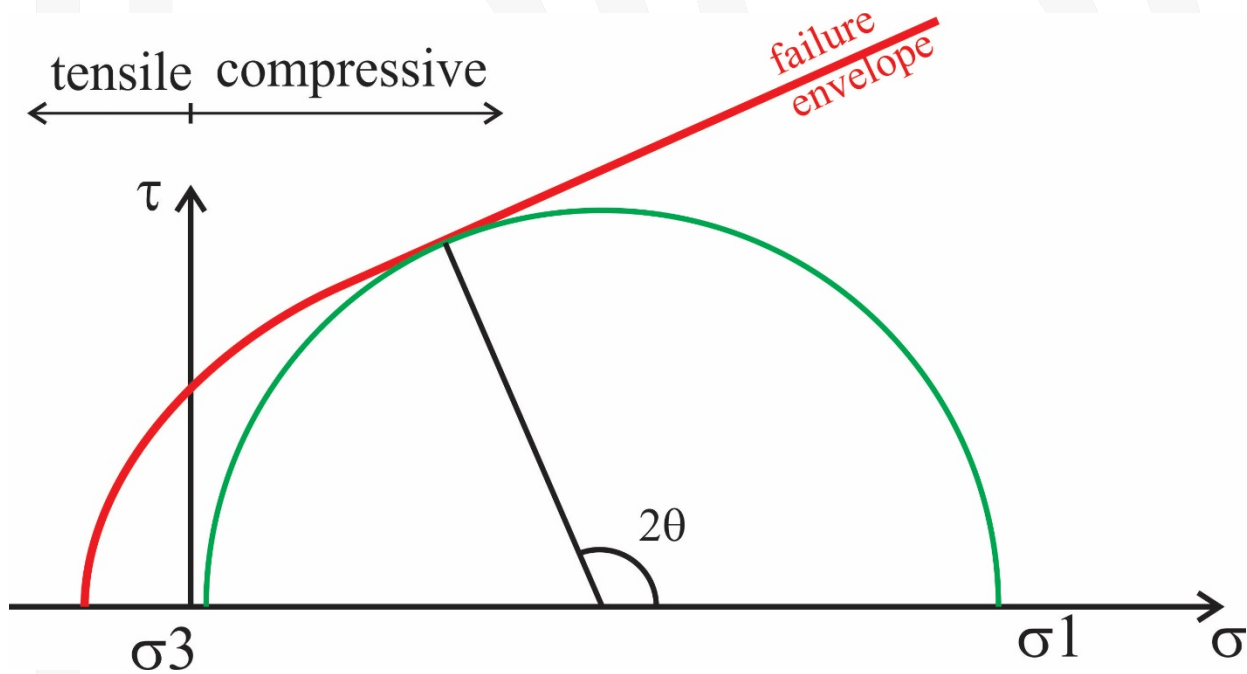
Fractures in a sandstone-dominant unit



Fractures in a shale unit



Mohr's circle: Reminder



It's a convenient and geometrical approach to represent the stress state

It links the normal stress (σ) to the shear stress (τ)

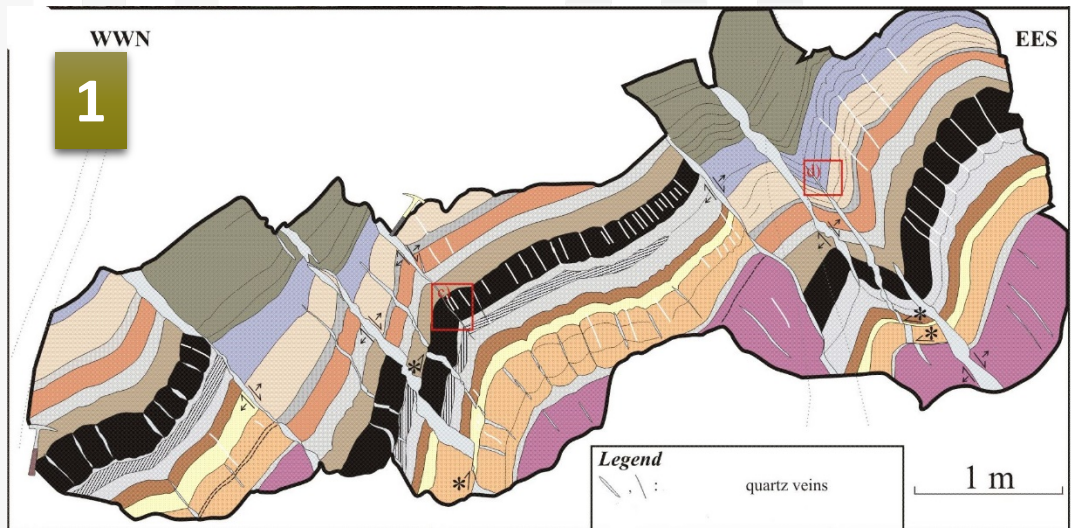
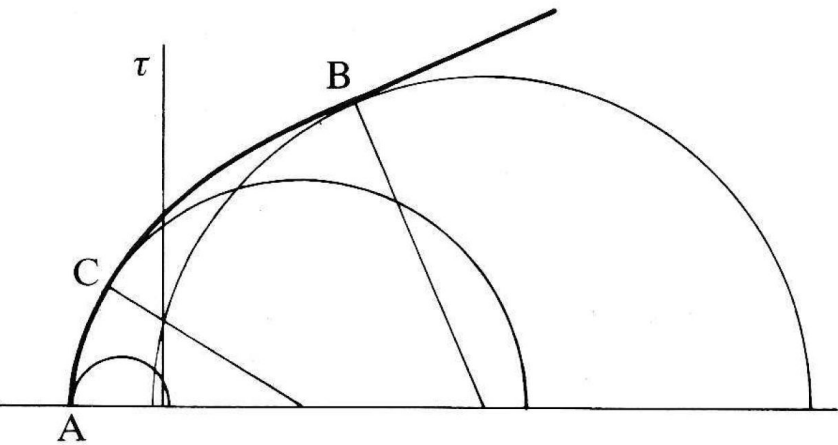
The failure envelope defines the differential stress ($\sigma_1 - \sigma_3$) required for a failure to occur

The newly-formed fracture develop at an angle θ with respect to σ_1 orientation

There are 2 main domains: tensile and compressive

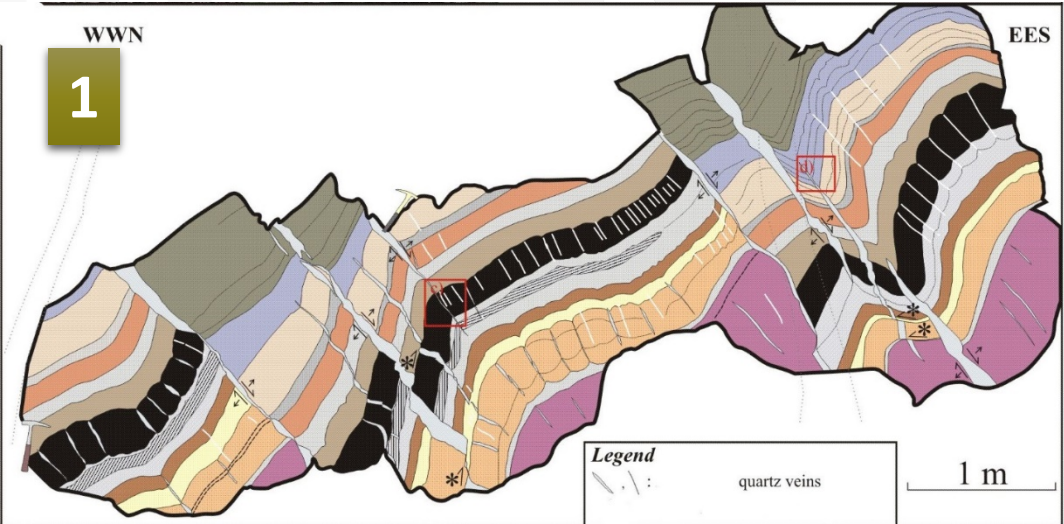
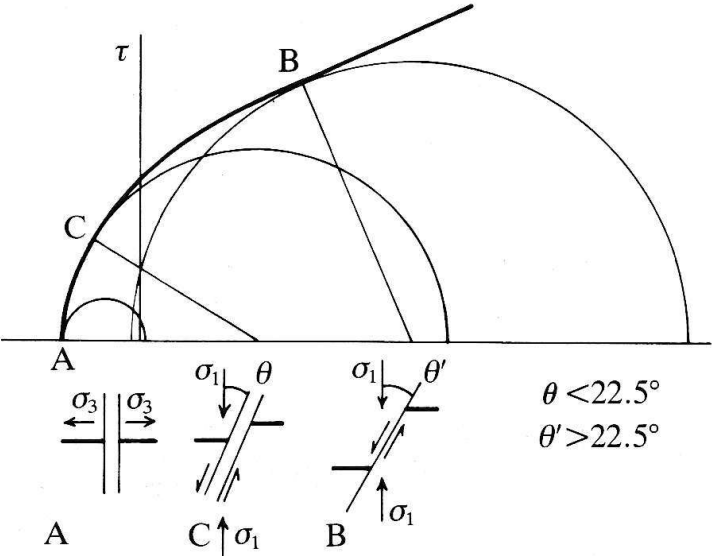
Tensile stress conditions can occur in rocks due to high fluid pressure

Exercise 2 – Classify the fractures according to their Morh's circle



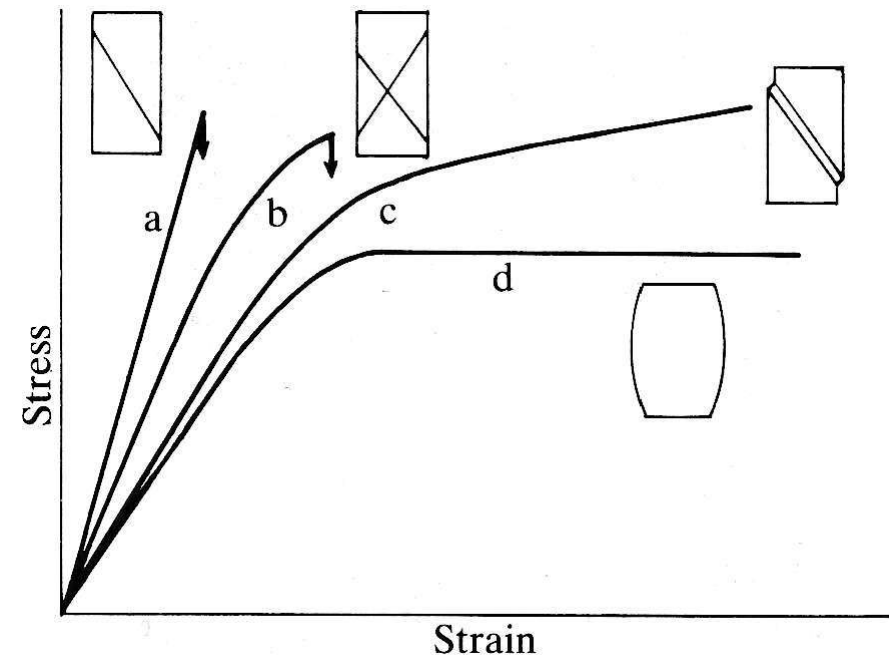
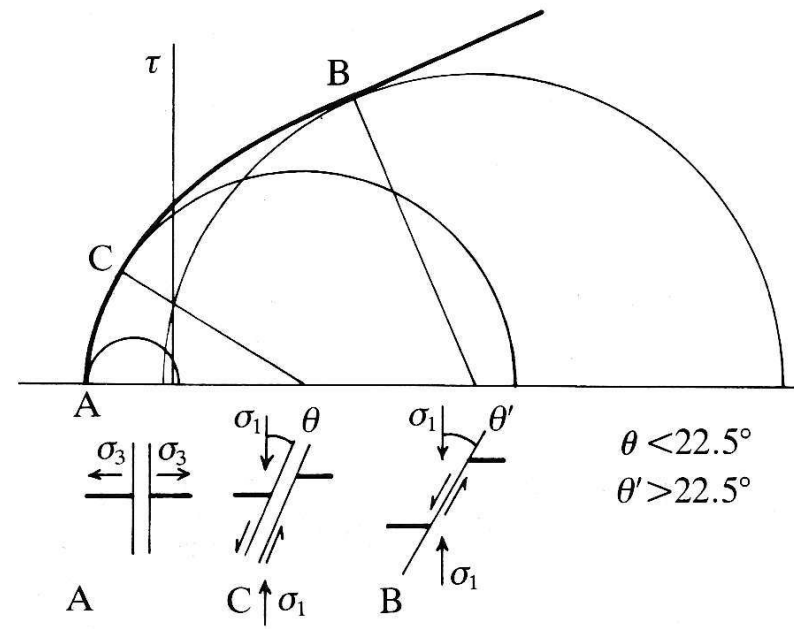
Circle-A	Circle-B	Circle-C

Exercise 2 – Classify the fractures according to their Morh's circle



Circle-A	Circle-B	Circle-C
2	3	1

Rheological contrast





Interlayer of shale and
sandstone layers from the
Mardasson Quarry
(Bastogne, Belgium)



‘There seems to be one elephant left in the room that is still commonly overlooked or ignored in these numerical models: anisotropy.’ (Ran et al., 2018)

Thank you very much for your attention



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References



- Davis, D.S. Chapman, T.M. Van Wagoner & P.A. Armstrong (2007) Thermal conductivity anisotropy of metasedimentary and igneous rocks. JGR 112, doi:10.1029/2006JB004755
- Hao Ran, Tamara de Riese, Maria-Gema Llorens, Melanie A. Finch, Lynn A. Evans, Enrique Gomez-Rivas, Albert Griera, Mark W. Jessell, Ricardo A. Lebensohn, Sandra Piazzolo, Paul D. Bons (2018). Time for anisotropy: The significance of mechanical anisotropy for the development of deformation structures. J. Struc. Geol. DOI:10.1016/j.jsg.2018.04.019.
- Hartmann, A., Rath, V., Clauser, C., 2005. Thermal Conductivity from Core and Well Log Data, Int. J. Rock Mechanics and Mining Sciences, 42, pp. 1042-1055.
- H.B. Lynn, 2018. The Fabric, or Internal Structure, of Rocks - The patterns of Anisotropy. In: AAPG Explorer, Feb & March 2018. Satinder Chopra (Eds).
- Popov, Y., Bayuk, I., Parshin, A., Miklashevskiy, D., Novikov, S., Chekhonin, E., 2012. New methods and instruments for determination of reservoir thermal properties. Thirty-Seventh Workshop on Geothermal Reservoir Engineering, Stanford University, Stanford, California, January 30 - February 1, 2012. SGP-TR-194.
- Popov, Y., Pribnow, D.F.C., Sass, J.H., Williams, C., Burkhardt, H., 1999. Characterization of rock thermal conductivity by high-resolution optical scanning. Geothermics 28, pp 253-276.
- N.J. Price and J.W. Cosgrove (1990). Analysis of Geological Structures. Cambridge University Press, 502 p.
- R. Skomski (2008). Simple Models of Magnetism. Oxford University Press. DOI:10.1093/acprof:oso/9780198570752.001.0001
- D.L. Turcotte & G. Schubert (2002). Geodynamics: 2nd Edition, Cambridge University Press, 456 p.